

2 Supplementary Information for

3 Mean Precipitation Change from a Deepening Troposphere

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7 This PDF file includes:

8 Supplementary text

9 Figs. S1 to S12

10 References for SI reference citations

11 **Supporting Information Text**

12 **1. CRM Simulations**

13 **A. Organization.** The spontaneous organization or ‘self-aggregation’ of convection has been much studied recently (see the
14 review by 1). DAM, however, has not been shown to exhibit this behavior; indeed, the simulations of (2) were initialized in an
15 aggregated state precisely because DAM would not spontaneously aggregate. The simulations in this study are no different, as
16 shown in Fig. S1 below, which plots snapshots of column relative humidity (CRH) on the last day of each simulation. CRH
17 here is defined as the water vapor path $\int \rho_v dz$ (kg/m²) divided by its saturation value. No organization is evident, and the
18 low CRH values associated with aggregation (0.3 and below, see Fig. 6 of 3) are not observed here. Note that the absence in
19 DAM of both self-aggregation as well as a sub-grid turbulence scheme is consonant with the results of (4), who show that
20 entrainment of dry air into cloud updrafts via sub-grid turbulence parameterizations can be critical for aggregation.

21 **B. LW and SW flux divergence profiles.** The main text argues that $(-\partial_T F^{\text{LW}})(T)$ and $(-\partial_T F^{\text{SW}})(T)$ are separately T_s -invariant.
22 We confirm this in Figs. S2 and S3, which as in Fig. 2 plot $-\partial_T F$ profiles in z , p , and T coordinates, but for the LW and SW
23 bands separately.

24 **C. CRM clear-sky flux divergence profiles.** The argument given in the main text for the T_s -invariance of $-\partial_T F$ is a clear-sky
25 argument, but all-sky flux divergences are shown in Figs. 2, S2 and S3. We argue in the main text that this is permissible
26 because cloud fraction in these simulations never surpasses $\sim 10\%$ at any height, so it is the clear-sky physics which dominates.
27 This claim is supported by the left and center panels of Fig. S4, which shows that the clear-sky flux divergence profiles are
28 almost indistinguishable from the all-sky profiles in Figs. S2 and S3, and are also indeed T_s -invariant. The right panel of
29 Fig. S4 directly contrasts the all-sky and clear-sky $-\partial_T F^{\text{net}}$ profiles for the $T_s = 300$ K simulation, and confirms that the
30 cloud-radiative effect in these simulations is not dramatic.

31 **2. Optical depth profiles**

32 The main text argues that water vapor optical depth $\tau_\lambda(T)$ is T_s -invariant. This argument was put forth by (5) and (6),
33 but has to our knowledge never been explicitly checked with a comprehensive radiative transfer calculation. Doing so with
34 RRTM is not straightforward, however, as RRTM is a ‘correlated- k ’ model producing band-averaged output, where each band
35 (there are 16 in the LW) covers a wide range of absorption coefficients and optical depths (7). We thus turn to a different,
36 line-by-line radiative transfer model, RFM (8). Feeding average p , T , and specific humidity profiles into RFM with the water
37 vapor continuum turned on and no CO₂ produces the optical depth profiles shown in Fig. S5. These show a reasonable degree
38 of T_s -invariance across a wide range of surface optical depths (and hence absorption coefficients). Deviations from perfect
39 T_s -invariance are likely due to pressure broadening as well as changes in lapse rate $\Gamma(T)$ between simulations, but this requires
40 further investigation. Temperature scaling factors should not contribute to deviations from T_s -invariance since these are also
41 T_s -invariant functions of T (e.g. Eq. (4.62) of reference 9).

42 **3. GCM analysis**

43 **A. Variance of $-\partial_T F^{\text{net}}$ and $\Gamma(T)$.** Figure S6 plots the variance $\text{Var}(\Gamma)$ of $\Gamma(T)$ within T_s bins for various T_s for the IPSL
44 model (other models show similar results). A pickup in variance in the lower atmosphere is evident, and a candidate T_{ext} is
45 given by the minimum temperature satisfying $T > 240$ K (to avoid the large variance regions in the upper atmosphere) and
46 $\text{Var}(\Gamma) > 0.5 \text{ K}^2/\text{km}^2$, plotted in black dots and the dashed lines.

47 By Eqns. (4–6) of the main text this implies a similar pickup in variance in $-\partial_T F^{\text{net}}$, shown in Figure S7 (other models
48 again show similar results, and calculations using clear-sky fluxes show a similar sharp pickup in variance, though the relatively
49 large variances ultimately reached in the surface-based layers are sometimes smaller). We then obtain a second candidate T_{ext}
50 as the minimum temperature level satisfying $T > 240$ K and $\text{Var}(-\partial_T F^{\text{net}}) > 5 \text{ (W/m}^2/\text{K})^2$. This T_{ext} is again shown by
51 black dots and dashed lines, and values are reasonably close to those obtained from $\text{Var}(\Gamma)$. If the T_{ext} candidate derived from
52 $\text{Var}(-\partial_T F^{\text{net}})$ exists then it is used for T_{ext} , as it better represents where the AMIP and AMIP4K $-\partial_T F^{\text{net}}$ profiles diverge; if
53 this T_{ext} candidate does not exist (as for the $T_s=250$ K bin of the IPSL model), then T_{ext} as diagnosed from $\text{Var}(\Gamma)$ is used.

54 **B. AMIP_{ext} profiles for other T_s bins.** Figure 6 suggests that AMIP_{ext} profiles are often a good approximation to the AMIP4K
55 profiles, but this is not always the case (e.g. the IPSL panel). For a better sense of the robustness of agreement between
56 AMIP_{ext} and AMIP4K profiles, we show the analogous panels but for the $T_s=280$ K bins, rather than $T_s = 290$ K, in Fig. S8.
57 These show that AMIP_{ext} profiles are typically a good approximation to the AMIP4K profiles, and that a failure of these
58 profiles to line up seems to be the exception rather than the rule.

59 **C. GCM clear-sky flux divergence and relative humidity profiles.** In the main text we claimed that the near-surface features in
60 the GCM $-\partial_T F^{\text{net}}$ profiles in Figs. 5 and 6 were sometimes, but not always, due to cloud radiative effects (CRE). Figure S9
61 show both all-sky and clear-sky $-\partial_T F^{\text{net}}$ profiles for the AMIP case for all models for the $T_s = 270$ K bin, for which many
62 models show a significant near-surface CRE. Figure S10, which is analogous to Fig. S9 but for the $T_s = 290$ K bin, shows on
63 the other hand that in this T_s bin the near-surface CRE across models is less consistent and less significant.

64 Figure S11 supports the claim in the main text that T_s -binned RH profiles also exhibit T_s -invariance aloft, but have
 65 near-surface features which shift downwards with warming. RH profiles are binned exactly as for the radiative fluxes, as
 66 described in *Materials and Methods*.

67 Figure S12 shows all-sky $-\partial_T F^{\text{LW}}$ and $-\partial_T F^{\text{SW}}$ for the $T_s=290$ K (AMIP) and $T_s=294$ K (AMIP4K) bins for all our
 68 CFMIP models, demonstrating that the T_s -invariance in GCMs holds for both the LW and SW separately, just as for the CRM
 69 (Figs. S2 and S3).

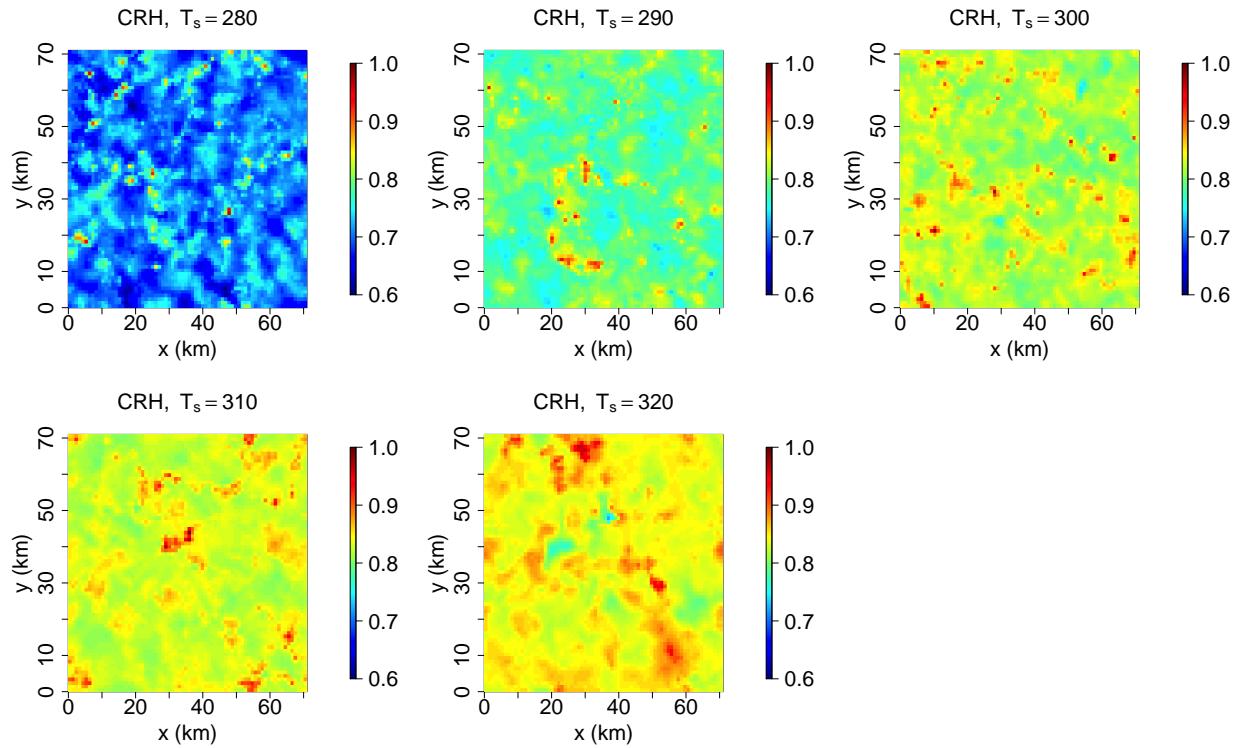


Fig. S1. Snapshots of column relative humidity (CRH) from the last day of each RCE simulation. No organization is evident, and the low CRH values associated with aggregation ($\lesssim 0.3$) are not observed.

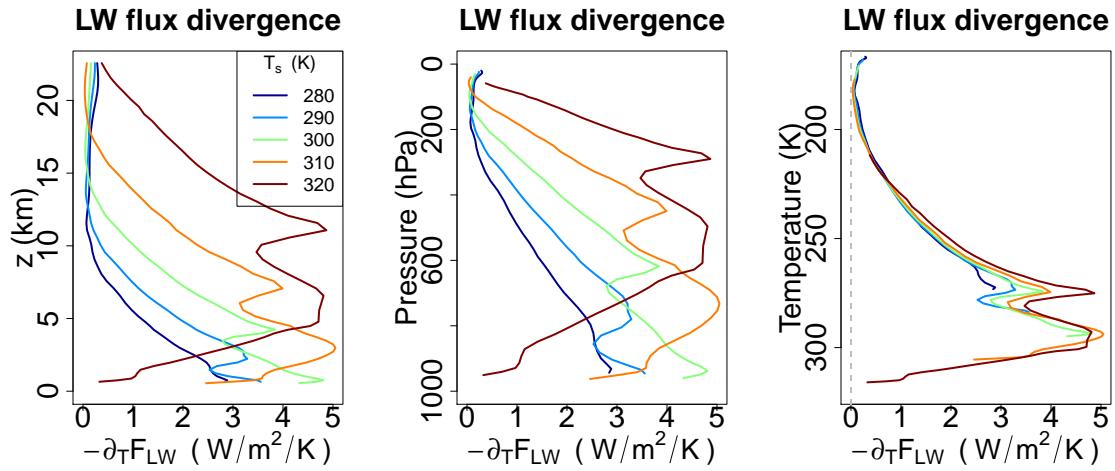


Fig. S2. LW flux divergence $-\partial_T F^{\text{LW}}$, as diagnosed from RRTM coupled to our CRM RCE simulations at $T_s=(280, 290, 300, 310, 320)$ K. Fluxes are plotted from the lifting condensation level of each simulation to 22.5 km for clarity, and in height, pressure, and temperature coordinates to emphasize the T_s -invariance of $(-\partial_T F^{\text{LW}})(T)$.

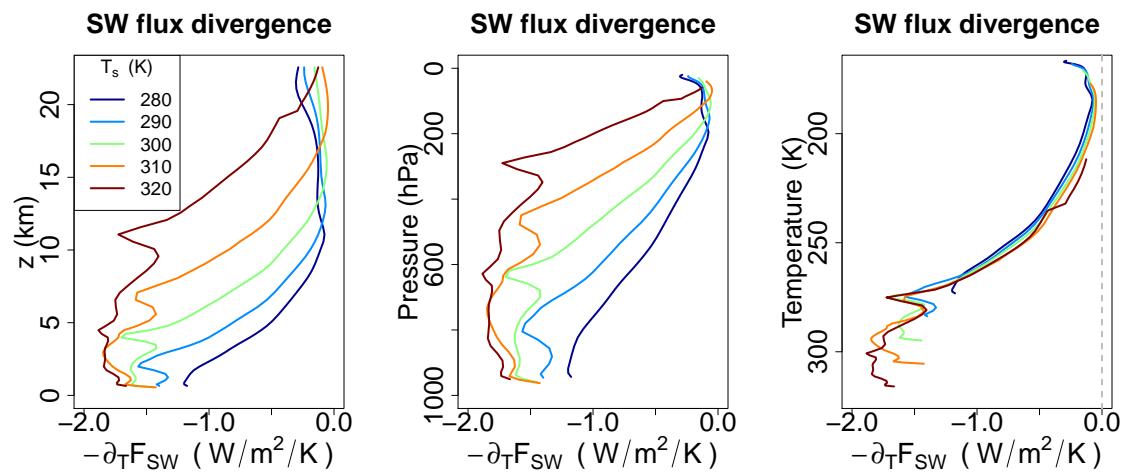


Fig. S3. As in Fig. S2, but for the SW band.

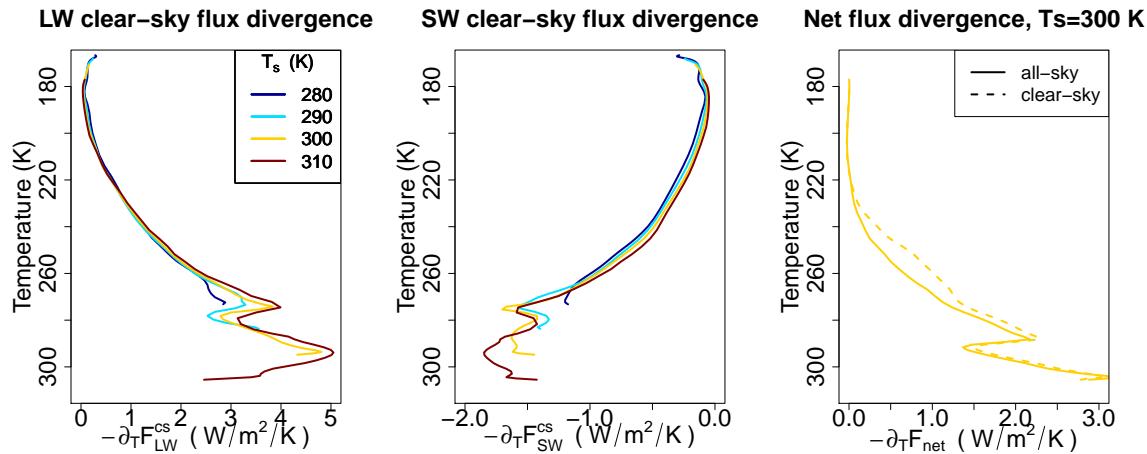


Fig. S4. **Left:** Clear-sky LW flux divergence $-\partial_T F_{\text{cs}}^{\text{LW}}$ **Center:** Clear-sky SW flux divergence $-\partial_T F_{\text{cs}}^{\text{SW}}$ **Right:** Clear-sky and all-sky net flux divergence for the $T_s = 300$ K simulations, all plotted as in Fig. 2. The left and center panels are almost identical to the right panels of Figs. S2 and S3, and the right panel above shows directly the small difference between the all-sky and clear-sky flux divergences for the $T_s = 300$ K simulation.

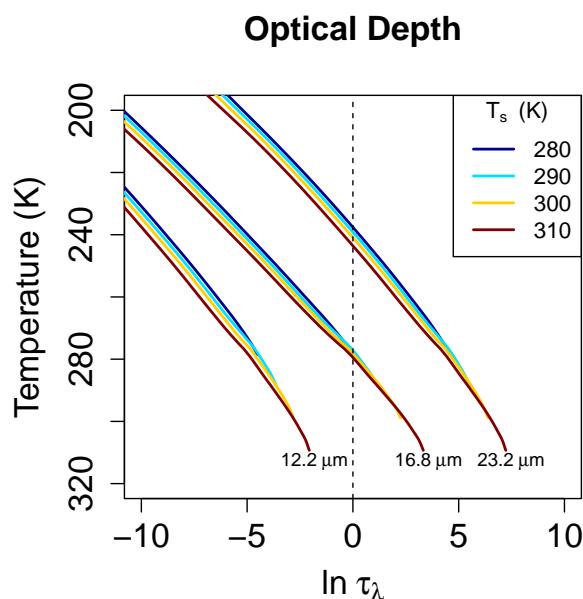


Fig. S5. Optical depth profiles $\tau_\lambda(T)$, obtained by feeding thermodynamic profiles from the RCE simulations into the RFM line-by-line radiative transfer code. Profiles are shown for water vapor only at three different wavelengths corresponding to surface optical depths of 0.01, 1, and 100 in the $T_s=280$ K simulation. A reasonable degree of T_s -invariance is seen at each wavelength: over the 30 K range of T_s , the temperature at which these lines reach $\tau_\lambda = 1$ for example (where cooling-to-space is maximized) varies by at most 8 K, or a little over 25% of the T_s range.

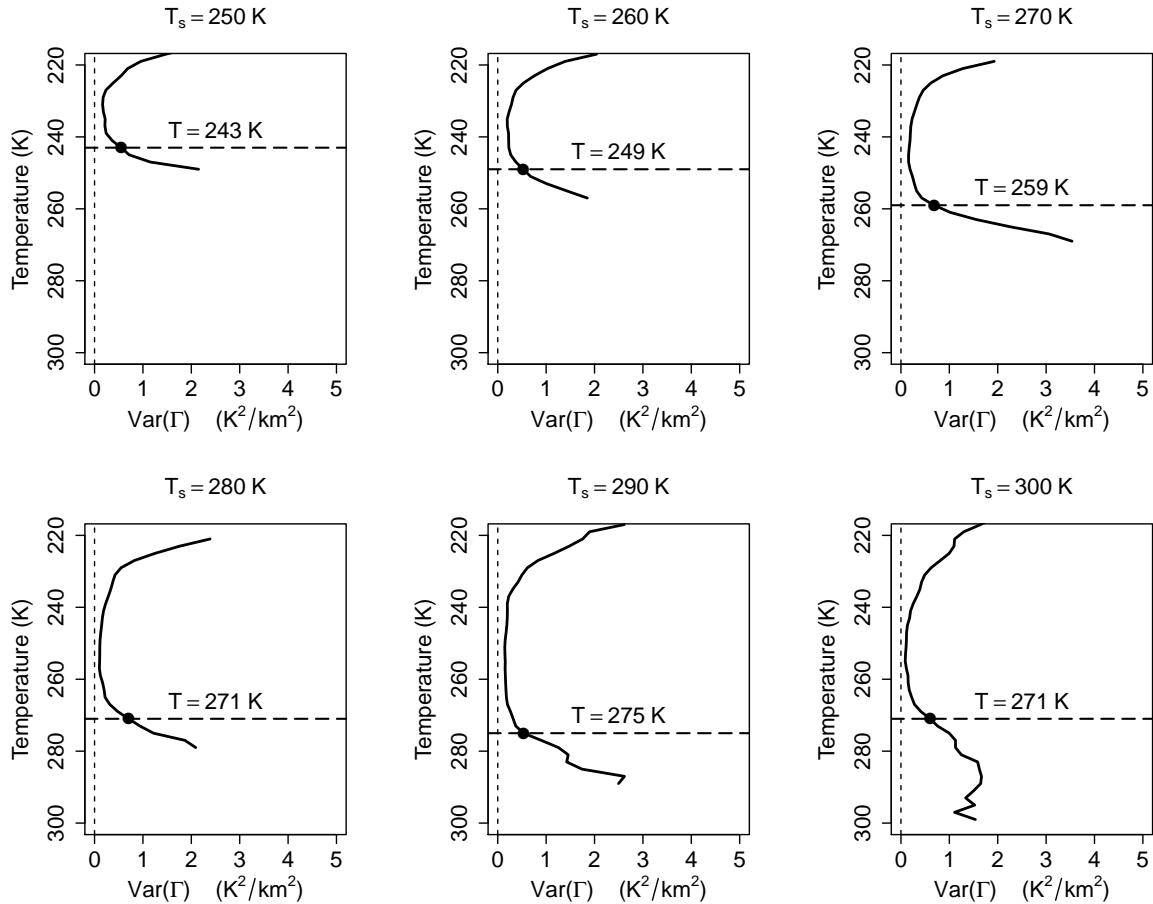


Fig. S6. Variance of $\Gamma(T)$ within T_s bins for the IPSL model. A fairly sharp pickup in the lower atmosphere is evident, similar to that found for $-\partial_T F^{\text{net}}$ profiles (Fig. S7). Black dots and dashed lines mark where the profiles exceed a threshold of $0.5 \text{ K}^2/\text{km}^2$ in the lower troposphere, showing that T where the variance in Γ picks up is comparable to the T_{ext} diagnosed from the variance in $-\partial_T F^{\text{net}}$.

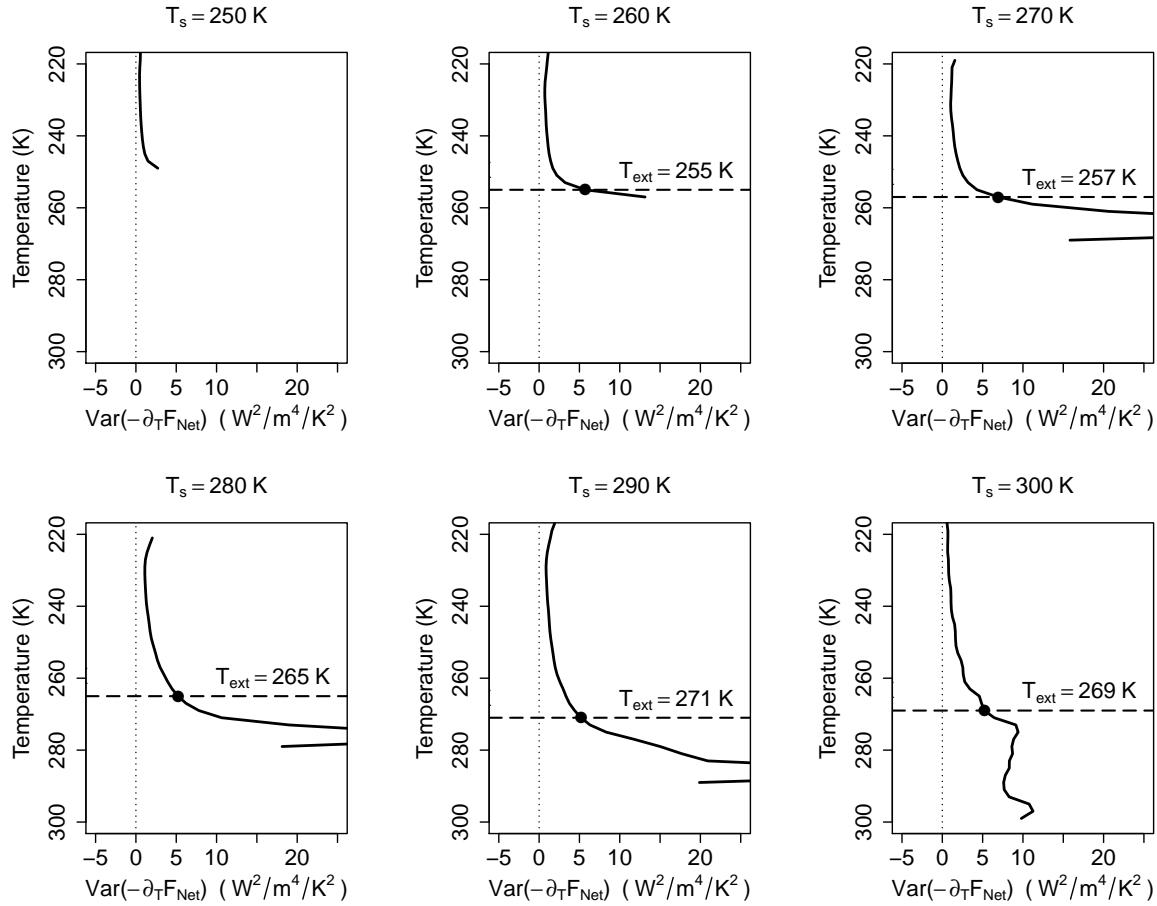


Fig. S7. Variance of $-\partial_T F^{\text{net}}$ within T_s bins for the IPSL model. A fairly sharp pickup in the lower atmosphere is evident for most bins, which is then used to diagnose T_{ext} , plotted in black dots. See SI text 3A for details.

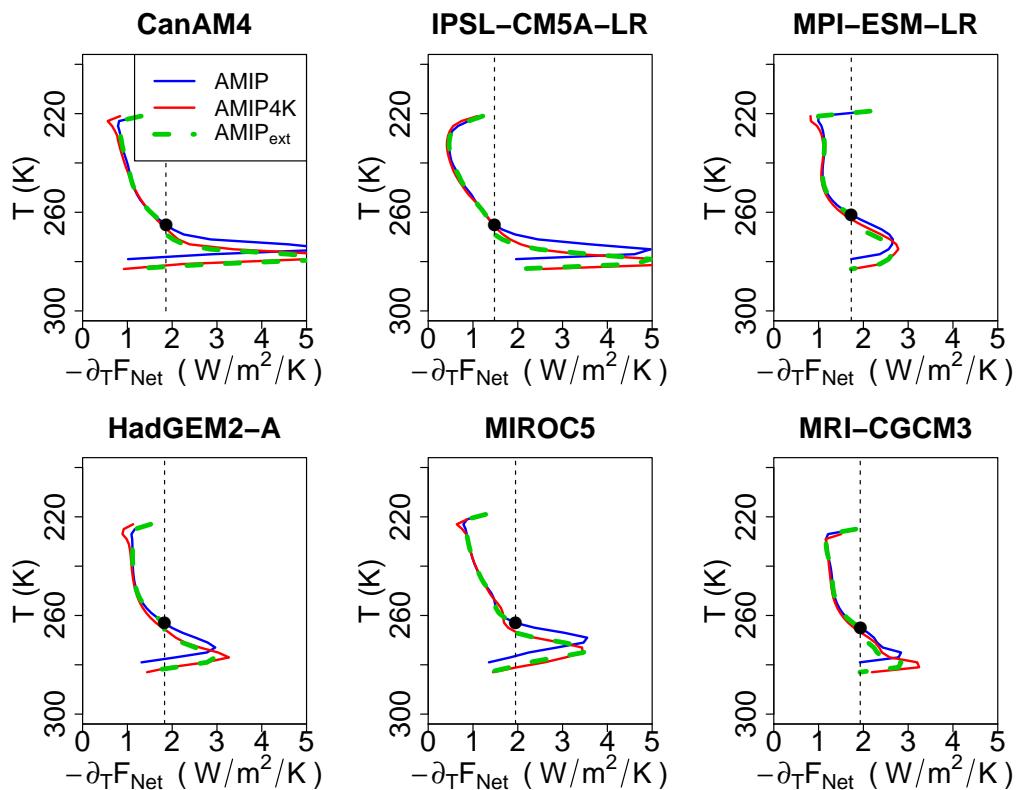


Fig. S8. As in Fig. 6 of the main text but for the $T_s = 280 \text{ K}$ bins.

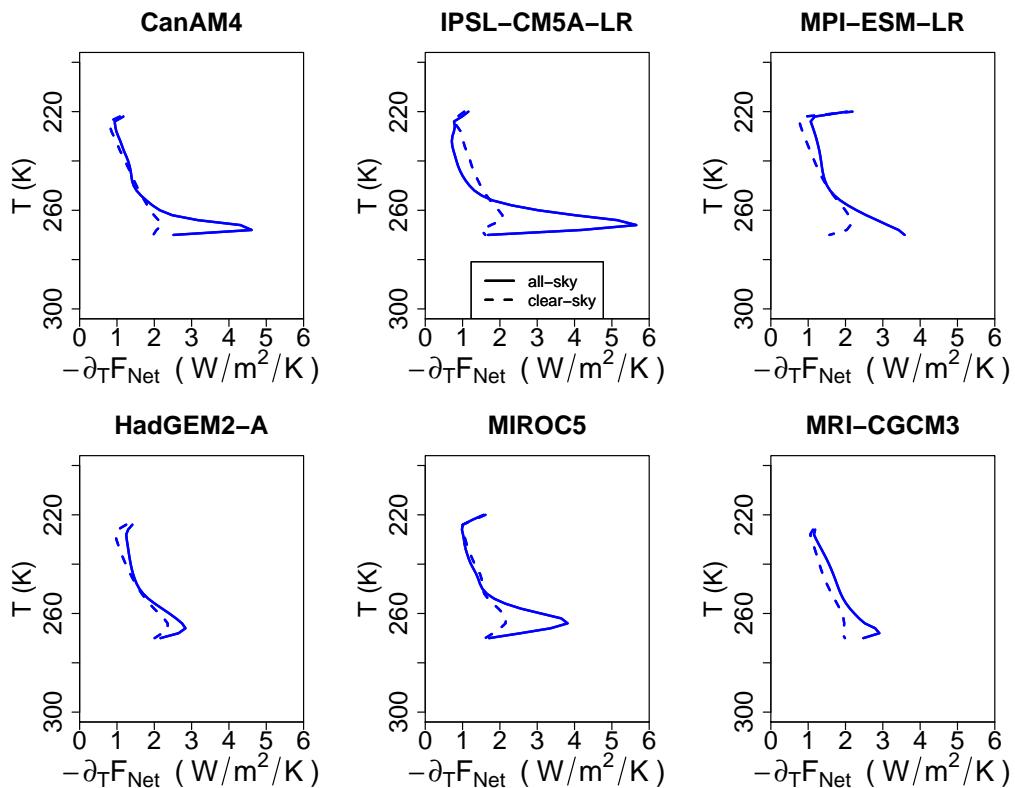


Fig. S9. All-sky and clear-sky $-\partial_T F^{\text{net}}$ profiles for the AMIP case for all models for the $T_s = 270$ K bin. The majority of models show a significant near-surface CRE.

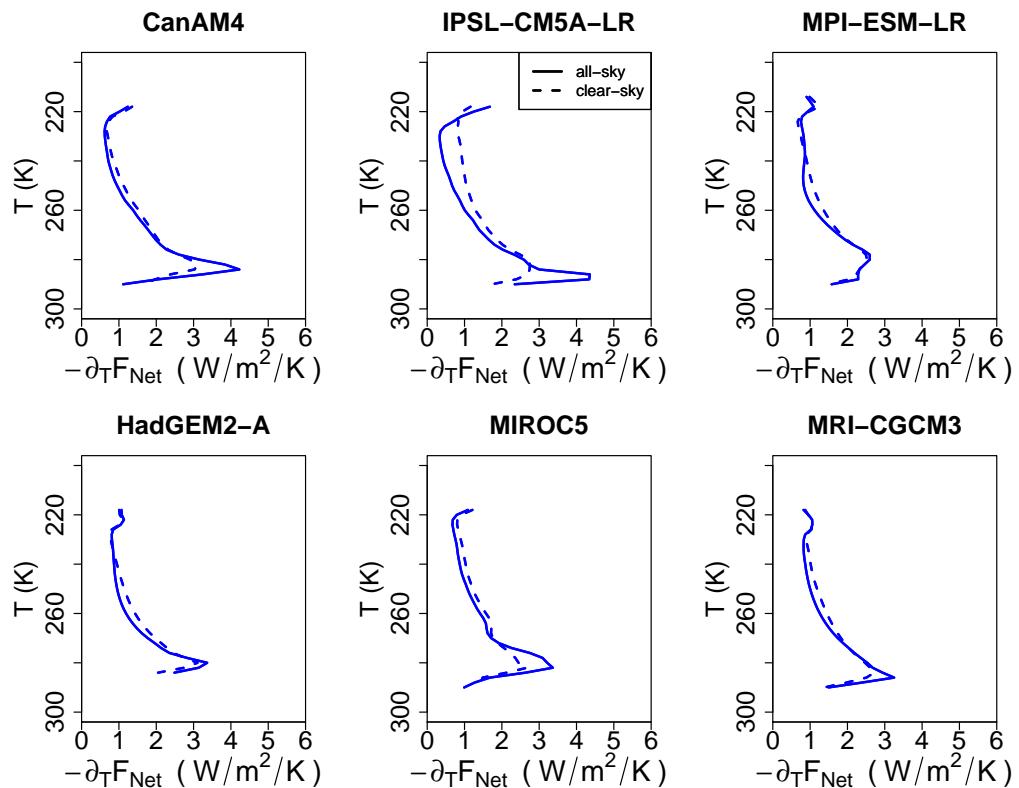


Fig. S10. As in Fig. S9, but for the $T_s = 290$ K bin. The near-surface CRE is much less significant across models than for the $T_s = 270$ K bin.

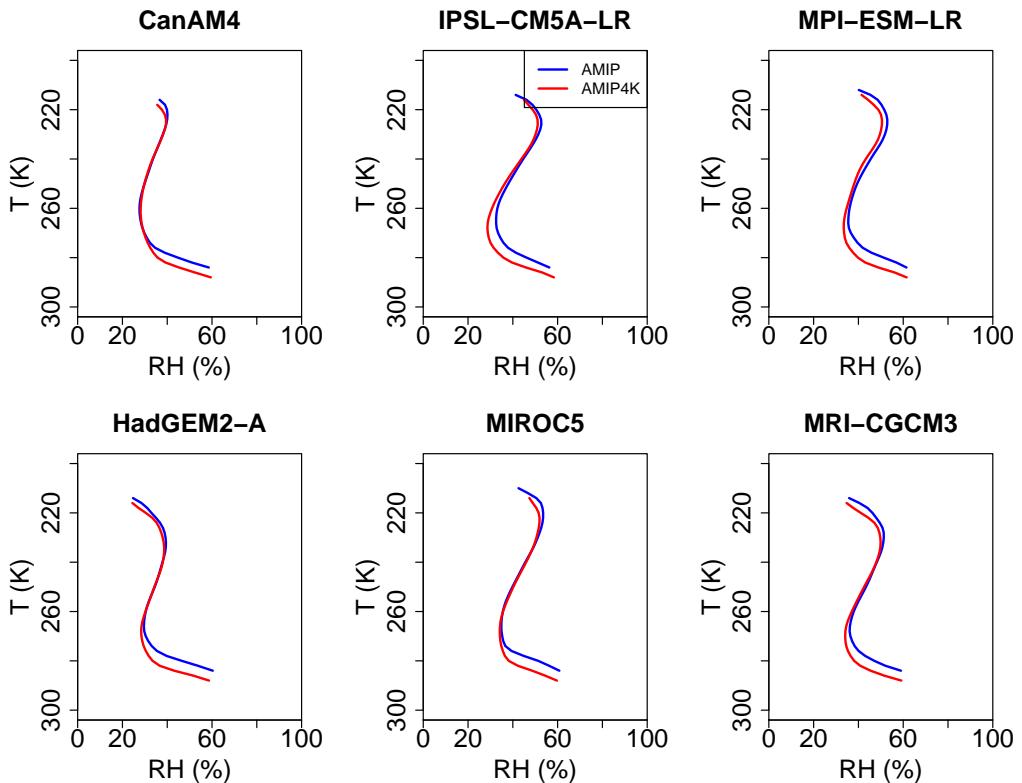


Fig. S11. RH profiles for our CFMIP models for the $T_s = 290$ (AMIP) and $T_s = 294$ K (AMIP4K) bins, computed just as for radiative fluxes. Like the $-\partial_T F^{\text{net}}$ profiles, the RH profiles show T_s -invariance aloft, but have lower-tropospheric features which shift downward (in temperature space) with warming.

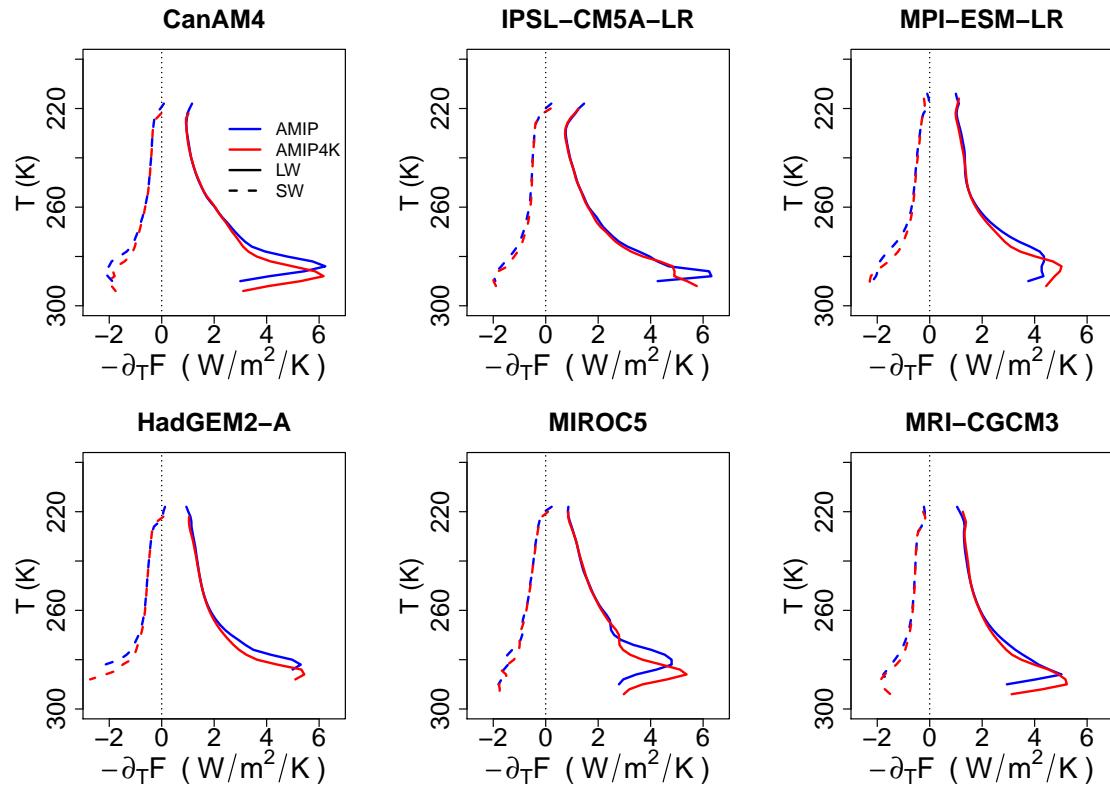


Fig. S12. Profiles of $-\partial_T F^{\text{SW}}$ and $-\partial_T F^{\text{LW}}$ for the $T_s=290 \text{ K}$ (AMIP) and 294 K (AMIP4K) bins for all six CMIP models. This is similar to Fig. 6 of the main text, but with the flux divergence decomposed into the LW and SW bands to show that T_s -invariance in the mid and upper troposphere holds across models holds for both the LW and SW separately, just as for the CRM.

70 **References**

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